Towards a System for Modeling the Impact of Ocean Acidification on Sea Scallops (*Placopecten magellanicus*) in Massachusetts Bay

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Abstract

The world’s oceans serve a critical role in absorbing CO₂ emissions, but this process contributes to rising ocean acidity. Northern coastal bodies of water like Massachusetts Bay are especially susceptible to ocean acidification. Massachusetts Bay supports a strong mollusk fishery, including the economically important Atlantic sea scallop (*Placopecten magellanicus*). As oceans acidify, the saturation state of aragonite, a mineral critical in mollusk shell formation, decreases, impeding shellfish development. In this study, a program was designed to evaluate the impact of ocean acidification on sea scallop growth in Massachusetts Bay. Equations modeling shellfish growth from previous literature were analyzed, and relevant formulae extracted and compiled. The resultant program was coupled to an existing MIT-MSEAS PE model. Initial parameters were sourced from recent population assessments. A total of four experimental scenarios were run for 32.5 days, each scenario depicting a varying dependency on three variables: aragonite availability, temperature, and physical phenomena (e.g., tides, winds, solitons, density-driven flow, internal waves). Growth of Massachusetts Bay sea scallops was found to depend heavily on aragonite saturation. Examined in isolation, variations in initial temperature did not substantially affect sea scallop growth. The incorporation of physical phenomena, however, resulted in subtle changes in scallop growth, demonstrating that ocean dynamics play an important role in temperature redistribution. Similarly, when both aragonite concentration and physical phenomena were incorporated into the model, the resultant effect on scallop growth indicated that ocean dynamics are a crucial driver of aragonite distribution as well. Coupling a shellfish growth submodel to the MIT-MSEAS PE allowed successful modeling of the impact of aragonite, temperature, and physical phenomena on sea scallop growth in Massachusetts Bay.

Summary

Levels of atmospheric carbon dioxide have increased as a result of human activity, causing oceans to acidify. Northern coastal waters like Massachusetts Bay are especially vulnerable to ocean acidification. As oceans acidify, the volume of minerals used in mollusk shell formation (e.g., aragonite) decreases. This study designed a computer program modeling the impact of ocean acidification on shellfish growth in Massachusetts Bay. Due to their vital economic importance, sea scallops were selected as the study species. Four scenarios were run, each with different variable dependencies. The results show that sea scallops rely heavily on aragonite saturation to grow and develop. The effects of temperature appear less significant and nearly uniform across locations. Physical dynamics are found to be critical in transporting aragonite to deepwater environments and considerably impact the growth of sea scallops.
1 Introduction

Worldwide, levels of atmospheric carbon dioxide (CO$_2$) continue to rise dramatically. Anthropogenic CO$_2$ emissions are of special concern, as they have exhibited an exponential pattern of growth since measurements were first taken at the Mauna Loa Observatory in 1958 [1]. Recent estimates by Friedlingstein et al. [2] place total anthropogenic CO$_2$ emissions from 1850 to 2020 at close to 2420 gigatons of CO$_2$. Elevated levels of atmospheric CO$_2$ trap more radiation, warming the earth and damaging the biosphere [3, 4]. If unaddressed, these emissions can precipitate natural disasters, scarcity of natural resources, and political discord [5].

Acting as carbon sinks, the world’s oceans are critical in countering elevated levels of CO$_2$ in the atmosphere. Since the beginning of the industrial era, oceans have absorbed almost half (48%) of all anthropogenic CO$_2$ emissions [6]. This buffer process releases hydrogen protons (H$^+$), decreasing ocean pH levels and increasing ocean acidity [7]. Not all marine zones are equally affected by ocean acidification. Coastal waters are particularly vulnerable, as elevated rates of freshwater input negatively impact buffering capability [8]. Northern waters exhibit lower buffering capacity than their southern counterparts due to lower average temperatures. Combining both of these features, northern New England coastal waters (e.g., Massachusetts Bay) are therefore especially imperiled by ocean acidification [9].

As oceanic pH levels decrease, the availability of calcium carbonate polymorphs like aragonite and calcite (minerals critical for shellfish growth and development) also decreases [10]. Of special concern is the decrease in aragonite saturation (Ω) due to ocean acidification. Aragonite is a more soluble form of calcium carbonate than calcite and is relied upon by shellfish larvae for initial shell formation [11]. The heightened solubility of aragonite, however, makes it susceptible to ocean acidification, which corrodes aragonite [12]. A decrease in the saturation state of aragonite negatively impacts shellfish growth, initiating a decline
in shellfish population size [12]. Ecologically, shellfish serve as critical building blocks of the complex oceanic trophic system. Acting as “ecosystem engineers”, shellfish filter water, combat erosion, and create habitable environments for other species [13]. Due to the durability of mollusk shells, this role persists long after shellfish death [14]. Shellfish contribute significantly to the global economy, but projections show that ocean acidification may lead to a $100 billion loss in shellfish-generated revenue by the year 2100 [15].

Of particular relevance to the Atlantic coast is the Atlantic sea scallop (*Placopecten magellanicus*). This species is of regional economic importance, with the 2019 commercial sea scallop harvest generating over $570 million in landings [16]. Although primarily concentrated on Georges Bank and the Mid-Atlantic Bight, sea scallops are distributed as far south as North Carolina and as far north as Newfoundland [17]. In the acidification-prone waters of Massachusetts Bay, a population exists along Stellwagen Bank [18]. In order to protect this ecologically and economically critical species, the impact of ocean acidification on sea scallops must be examined.

This study is designed to demonstrate the relationship between ocean acidification and sea scallop population distribution. Values for ocean acidification variables are first generated by a primitive equation model, then transferred to a scallop growth model. In this study, initial ocean acidification variables (aragonite saturation and temperature) are sourced from a numerical simulation of Massachusetts Bay using the Massachusetts Institute of Technology Multidisciplinary Simulation, Estimation, and Assimilation Systems primitive equation (MIT-MSEAS PE) model. To design an applicable shellfish growth model, existing models were reviewed and the roles of ocean acidification variables in shellfish growth and development were determined. Several shellfish models have been produced previously, but with varied emphasis (e.g., individual growth, scope for growth, population dynamics, fisheries management) [19] [20] [21]. After extensive evaluation, an existing scallop growth model outlined by Cooley et al. [21] in 2015 was selected and adapted to Massachusetts Bay conditions.
In this way, a set of representative equations reflecting shellfish growth was established and coupled to the existing MIT-MSEAS PE model. The resultant program serves as a comprehensive, real-time detection system displaying the impact of ocean acidification on sea scallops in Massachusetts Bay.

2 Methods

2.1 Categorizing Existing Shellfish Growth Models

First current progress in shellfish growth modeling was reviewed. Available models had a wide range of applications, forcing variables, and conclusions [22, 23]. The pool of models was divided into two categories for further analysis: individual growth models and population dynamics models. Individual growth models were found to often rely on dynamic energy budget (DEB) theory. DEB theory is a metabolic theory that represents organism energy use, reproduction, and development as functions of size and environmental interactions [24]. In addition, DEB theory is general and easily applied to a wide range of species across developmental stages [25]. Models implementing DEB theory, however, did not reflect the influence of established ocean acidification variables [26]. In contrast, population dynamics models incorporated a larger bank of equations for a detailed analysis of an entire population. Within these equations, especially those related to scallop growth and recruitment, variables influenced by ocean acidification were found to play primary roles [21].

2.2 Suitability of Model Categories for Project Goals

This study determined that the ideal method of measuring the impact of ocean acidification on Atlantic coast shellfish would utilize population dynamics models. Individual growth models, although more simple, would reveal less information on the overall state of a
species as compared to population dynamics models. Several models integrating temperature \[23, 27\] were identified. Only one (Cooley et al., 2015) \[21\] addressed a second relevant ocean acidification variable: aragonite. Following the process outlined by Cooley et al., the current study adapts and couples this model to the MIT-MSEAS PE model and is the first to apply it to sea scallops in the Massachusetts Bay area.

2.3 Shellfish Growth & Population Dynamics Equations

The shellfish submodel nested in the integrated assessment model designed by Cooley et al. was selected as most suitable for modeling sea scallop response to ocean acidification \[21\]. Unless otherwise stated, all equations below are sourced from the integrated assessment model outlined in the Cooley study.

According to the shellfish submodel, total scallop population size is comprised of several size class bins, which scale by 5 mm intervals. Scallops enter the first bin at the shell height \(H\) of 40 mm, which corresponds to the time that they recruit into the general sea scallop population \[28\]. This study defines the last size bin as having a lower limit of 90 mm and an upper limit of \(\infty\) to account for all sea scallops. The 90mm lower limit corresponds to the 3.5-inch minimum shell height regulation for commercially harvested sea scallops in Massachusetts, leaving open the future possibility of considering fisheries pressures upon this final bin \[29\].

The number of individuals in each bin comprise \(h\) elements that then make up the vector \(\vec{n}(t)\). The sum of all the elements of \(\vec{n}(t)\) represents the total sea sea scallop population at each time step \(t\) (yr).

The total count of scallops in each size class is given as:

\[
\vec{n}(t + \Delta t) = \vec{R} \Delta t + G_t \cdot e^{-(M)\Delta t} \cdot \vec{n}(t)
\]  

(1)
where $\vec{n}(t + \Delta t)$ is the vector representing the subsequent scallop population (millions) after the time step is applied, $\vec{R}$ is a measure of recruitment (individuals yr$^{-1}$), $\Delta t$ is the change in time (yr), $G_t$ is a matrix which denotes the probability of advancing bins, $M$ is a measure of natural mortality (yr$^{-1}$), and $\vec{n}(t)$ signifies the scallop population before the time step is applied.

The equation modeling recruitment is defined as:

$$\vec{R} = \frac{\alpha R SSB}{\gamma R + SSB}$$  \hspace{1cm} (2)

where $\alpha_R$ is defined as the recruitment asymptote (millions of individuals), $\gamma_R$ is defined as the half saturation coefficient (mT meats), and $SSB$ is the spawning stock biomass (mT meats). The values for $\alpha_R$ and $\gamma_R$ were obtained from the Georges Bank initial parameters outlined in Cooley, et al. [21]. Scallops only recruit into the first (40 – 45 mm) size class. For simplicity, this study assumes that the recruitment rate is constant throughout the year.

To model the impact of aragonite saturation state ($\Omega_{aragonite}$) on scallop recruitment, an empirical scale factor is computed to be applied to initial spawning stock biomass ($SSBi$):

$$SSB \text{ scale factor} = \frac{\% \text{ survival}}{\% \text{ average survival}} = \frac{(20.5\Omega_{aragonite} - 3.7)}{45}$$  \hspace{1cm} (3)

where $\% \text{ survival}$ is represented by the function $(20.5\Omega_{aragonite} - 3.7)$. The value for $\% \text{ average survival}$ is given as 45 in accordance with control survival rates of great scallops ($Pecten maximus$) established by Andersen et al. [30]. This scale factor is then applied to $SSBi$ following the relationship outlined in Equation 4:

$$\text{scaled } SSB = (SSBi) \text{ (SSB scale factor)}$$  \hspace{1cm} (4)
where scaled SSB is the output to be used in quantifying R, SSBi is initial spawning stock biomass, and SSB scale factor is the scale factor applied to initial spawning stock biomass. Calculated at each hour-long advance in time, scaled SSB serves as an input to Equation 2 which models sea scallop recruitment adjusted for aragonite saturation.

The growth matrix $G_t$ describes scallop growth across size classes at every active grid point. In general, $G_t$ is a square matrix of size $(n_{\text{bins}} \cdot n_{\text{gridpoints}}) \times (n_{\text{bins}} \cdot n_{\text{gridpoints}})$ where $n_{\text{bins}}$ is the total number of size classes being considered and $n_{\text{gridpoints}}$ is the number of active grid points (points where scallops are being modeled). This study’s program contained 13,139 active locations. $G_t$ becomes especially tractable when the following modeling assumptions are made.

1. Scallop growth at one grid point is not affected by scallop growth at any other grid point. This encompasses both the notion that scallops do not move frequently enough to travel between grid points and the idea that scallop growth does not impact the physical fields (temperature, aragonite) that are controlling growth. Under this assumption, $G_t$ becomes block diagonal with $n_{\text{gridpoints}}$ block sub-matrices, $\hat{G}_t$, of size $n_{\text{bins}} \times n_{\text{bins}}$.

2. The time step $\Delta t$ is small enough that scallops cannot “jump” size classes (e.g., scallops from size class $i - 1$ can grow into size class $i$ but scallops from size classes 1, 2, ..., $i - 2$ cannot). With this assumption each of the block sub-matrices, $\hat{G}_t$, becomes bi-diagonal with the nonzero elements restricted to the main diagonal and to the first sub-diagonal.

If the probability of a scallop growing out of size class $i$ is denoted as $P_i$, then each of the elements, $\hat{G}_{t;(i,j)}$, of $\hat{G}_t$ can be written as:

$$
\hat{G}_{t;(i,j)} = \begin{cases} 
1 - P_i & \text{if } j = i; \ i \in \{1, 2, \ldots, n_{\text{bins}}\} \\
N_i - 1 & \text{if } j = (i - 1); \ i \in \{2, 3, \ldots, n_{\text{bins}}\} \\
0 & \text{otherwise}
\end{cases}
$$
To evaluate $P_i$, the range of shell lengths that could grow out of size class $i$ is estimated by computing how much growth a shell beginning as the upper limit of size class $i$ can grow in time $\Delta t$ ($\Delta H$). Then, it is assumed that any scallop of shell length within $\Delta H$ of the upper limit will grow out of size class $i$. Finally, it is assumed that the scallop shell lengths within a size class are uniformly distributed in that size class, so the probability is ratio of the lengths:

$$P_i = \frac{H_{t+\Delta t} - H_t}{\text{Bin Width}} = \frac{\Delta H}{\text{Bin Width}}$$  \hfill (6)

where $H_{t+\Delta t}$ is shell length (mm) after the change in time, $H_t$ signifies shell length before the time step is applied, and $\text{Bin Width}$ represents the width of the size bin (5 mm).

Change in shell length $\Delta H$ is represented by:

$$\Delta H = (H_\infty - H_t) \left(1 - e^{-K \Delta t}\right)$$  \hfill (7)

where $H_\infty$ is asymptotic length (mm) and $K$ is the Brody growth coefficient, a measure of how fast it takes for an organism to reach maximum length.

$K$ is calibrated to both temperature and aragonite. To calculate change in $K$ because of change in aragonite saturation, it is first necessary to determine the relative change in aragonite from the control:

$$\Delta \Omega = \frac{\Omega}{\Omega_{\text{control}}} - 1$$  \hfill (8)

where $\Delta \Omega$ is the relative change in aragonite from the control, $\Omega$ is the measured aragonite saturation state, and $\Omega_{\text{control}}$ is the control aragonite saturation state.
Then, change in growth rate is represented by the empirical equation:

$$\Delta G = 1.272\Delta \Omega + 0.075$$ (9)

where $\Delta G$ is a measure of relative change in growth rate.

Finally, assuming that the relative change in $G$ is equal to the relative change in $K$, change in $K$ from the initial value is generated by multiplying relative change in growth rate by the Brody growth coefficient:

$$\Delta K_\Omega = \Delta G \cdot K_i$$ (10)

where $\Delta K_\Omega$ is the relative change in $K$ from the initial value $K_i$ due to change in aragonite saturation.

To calibrate $K$ to change in temperature, this study drew from empirical temperature-growth relationships outlined in Heilmayer et al. [31] and determined that the sum of $K_i$ and the change in $K$ due to temperature can be represented by the empirical equation:

$$K_i + \Delta K_T = \frac{e^{4.22 - 958.461 \cdot \frac{1}{T}}}{M_\infty}$$ (11)

where $K + \Delta K_T$ represents $K$ plus the change in $K$ due to temperature, $T$ is a measure of temperature (kelvin), and $M_\infty$ is the asymptotic total weight (g).

After the outputs of Equations 10 and 11 are calculated, overall change in $K$ due to change in $T$ and $\Omega$ can be summed:

$$K_{T,\Omega} = K_i + \Delta K_T + \Delta K_\Omega$$ (12)
where $K_{T, \Omega}$ represents change in $K$ due to both temperature and aragonite saturation.

### 2.4 Identification of Suitable Input Parameters

Several parameters from Cooley et al. [21] were adjusted. Population size (millions), length frequencies, and $SSBi$ (mT meats) were determined according to data collected by Lisi et al. in 2021 [18]. Asymptotic total weight $M_\infty$ was calculated following total weight to meat weight relationships outlined by Hennen and Hart [32]. A value for the control saturation state of aragonite ($\Omega_{\text{control}}$) was obtained by averaging the initial values for $\Omega$ at the bottom level around Stellwagen Bank. This study did not consider pressures exerted by fisheries (denoted as $Z_t$ in Cooley et al. [21]), allowing focused examination of the effect of ocean acidification on sea scallop performance. Initial values for dissolved inorganic carbon and total alkalinity were obtained from the MIT-MSEAS PE model and used to compute the aragonite saturation state using the CO2SYS equilibrium [33]. To best reproduce the environmental conditions of Massachusetts Bay, total dissolved inorganic silicon and total dissolved inorganic phosphorus were set at 0 $\mu$mol/kg. Standard atmospheric pressure was set at 10.132501 decibars. The $K_1$ and $K_2$ dissociation constants were set at 3 in accordance with the conclusions of Hannson refit by Dickson & Millero [34, 35]. The value for the KSO$_4$ dissociation constant was also set at 3 following the conclusions of Dickson in 1990 [36].

Table 1: Input Parameters for Sea Scallop Growth Model

<table>
<thead>
<tr>
<th>Parameter</th>
<th>Value</th>
<th>Units</th>
<th>Source</th>
<th>Use</th>
</tr>
</thead>
<tbody>
<tr>
<td>$\Omega_{\text{control}}$</td>
<td>1.75</td>
<td></td>
<td>MIT-MSEAS PE</td>
<td>Equation 8</td>
</tr>
<tr>
<td>average survival</td>
<td>45</td>
<td>%</td>
<td></td>
<td>Equation 3</td>
</tr>
<tr>
<td>$K_i$</td>
<td>0.429</td>
<td>yr$^{-1}$</td>
<td>21</td>
<td>Equation 10</td>
</tr>
<tr>
<td>$M_\infty$</td>
<td>467.685</td>
<td>grams</td>
<td>21, 32</td>
<td>Equation 11</td>
</tr>
<tr>
<td>$SSBi$</td>
<td>1013</td>
<td>mT meats</td>
<td>18</td>
<td>Equation 4</td>
</tr>
<tr>
<td>$\alpha_R$</td>
<td>20.17</td>
<td>millions</td>
<td>21, 37, 38</td>
<td>Equation 2</td>
</tr>
<tr>
<td>$\beta_R$</td>
<td>6.664</td>
<td>mT meats</td>
<td>21, 37, 38</td>
<td>Equation 2</td>
</tr>
<tr>
<td>$H_\infty$</td>
<td>144.9</td>
<td>millimeters</td>
<td>21, 37</td>
<td>Equation 7</td>
</tr>
<tr>
<td>$M_0$</td>
<td>0.12</td>
<td>yr$^{-1}$</td>
<td>21, 37</td>
<td>Equation 1</td>
</tr>
</tbody>
</table>
2.5 Program Design

The scallop growth program’s code was written in MATLAB, following a modular and flexible design approach. The program was designed from the top-down and written from the bottom-up. Construction began with the first equations necessary to calculate the variables involved in Equation 1 (e.g., Equation 3, Equation 8). Each function was unit tested along the build process. Substituting simple values for variables ensured that each individual file was performing as intended and that any subsequent errors would not be related to coding errors in functions already passing the unit tests. Values of aragonite saturation and temperature were passed to the functions through one-way coupling (without feedback) to the MIT-MSEAS PE model. The computational domain for the Massachusetts Bay simulation is shown in Figure 1.

Figure 1: Bathymetry chart of Massachusetts Bay providing visual representation of depth (meters) by location [39]. The modeling domain of the MIT-MSEAS PE model is enclosed by the red border.
2.6 Experimental Scenarios

Based on the initial parameters and equations outlined above, four experimental scenarios were run. The first scenario (Run 1) modeled scallop growth as a function of aragonite saturation, temperature, and physical phenomena (tides, winds, solitons, density-driven flow, internal waves). The second scenario (Run 2) modeled scallop growth as a function of temperature and physical phenomena, but not aragonite saturation. The third scenario (Run 3) modeled scallop growth as a function of temperature only, excluding both aragonite saturation and physical phenomena. The fourth and final scenario (Run 4) modeled scallop growth as a function of aragonite and temperature, but excluded physical phenomena. Each scenario was run from the time period August 11, 2019 to September 13, 2019 to ensure that accurate data from the MIT-MSEAS PE model was utilized.

2.7 Initialization

This study aimed to create initial conditions as representative of the Massachusetts Bay sea scallop population as possible. Although the shallow waters of the Bay hold mainly bay scallops, a sizeable population of sea scallops exists around Stellwagen Bank. Following the conclusions of the 2021 Stellwagen Bank sea scallop population assessment conducted by Lisi et al. [18], a set of area-specific initial conditions (population size, boundary parameters, and length distribution) were developed.

Stellwagen Bank sea scallops were divided into a northern and southern sector. The northern population box had latitudinal boundaries of 42.33° N to 42.62° N and longitudinal boundaries of 70.25° W to 70.67° W. The southern population box had latitudinal boundaries of 42.12° N to 42.33° N and longitudinal boundaries of 70.09° W to 70.44° W. Initial population sizes (89.65 million for the northern box and 17.76 million for the southern box) were assigned evenly within each box. Initial distribution of scallops across the 11 size bins
was calculated using size distribution data from Lisi et al. [18] for each box.

3 Results & Discussion

Within Massachusetts Bay, there exist several dynamical phenomena driving distribution of aragonite and temperature. Internal tides, winds, solitary waves, and background advection of water from the Gulf of Maine all play critical roles in influencing aragonite saturation state and temperature levels [39]. Deep waters such as those on the western border of Northern Stellwagen Bank (Figure 1, 42.3° N 70.5° W) tend to have lower saturation levels of aragonite than shallow waters (Figure 2a). This finding is due to acidic deep-water input from the Gulf of Maine [40], but local tidal and wave patterns result in mixing of aragonite across depths. While aragonite depth distribution without physics appears smoothly stratified (Figure 2a), oceanic disturbances from internal waves, etc. transport highly-saturated surface aragonite down to deeper waters (Figure 2b).

Figure 2: Effects of physical phenomena on aragonite distribution by depth over a 2.5-day period [39]. Note the disruption of aragonite stratification in Figure 2b due to the passing of an internal wave. These physical actions bring highly-saturated aragonite to deeper waters, making it more available for use by shellfish.
Figure 3: Run 1, change in sea scallop size distribution considering the effects of temperature, aragonite, and physics. Sea scallop growth shows a strong dependency on aragonite and physics. Figure generated by researcher.
Figure 4: Run 2, change in sea scallop size distribution considering the effects of temperature and physics, but not aragonite. Temperature influence appears fairly uniform, but subtle physical effects are present. Figure generated by researcher.
Figure 5: Run 3, change in sea scallop size distribution considering the effects of temperature, but not aragonite or physics. Note that the difference between Run 3 and Run 2 is small, displaying a slight influence of physics on temperature distribution and scallop growth. Figure generated by researcher.
Figure 6: Run 4, change in sea scallop size distribution considering the effects of aragonite and temperature, but not physics. The striking difference between Run 4 and Run 1 highlights the critical role ocean physics plays in aragonite distribution and scallop growth. Figure generated by researcher.
Over the 32.5-day period, the run with full dependency on aragonite, temperature, and physics (Run 1) saw substantial recruitment into the first (40 – 45 mm) size class (Figure 3a). The next two size classes (45 – 50 mm and 50 – 55 mm) exhibited increases in population size along the northern box depth decline, and the 50 – 55 mm size class saw growth along the western ledge of Northern Stellwagen Bank (Figure 3b, Figure 3c). The 55 – 60 mm size class appeared to have sizeable input from the lower size classes within an area of deep water on the western border of Northern Stellwagen Bank (Figure 3d). The southern zone, however, saw net loss in scallop population size. Mortality (M) is constant in the model. Therefore, an abnormal loss in population of a size class can either represent growth out of that size class into the class above, or that the number of scallops exiting the lower size class is less than the combined effect of growth into the upper size class and M. The following 6 size classes (60 – 90 mm) had similar patterns of growth (Figures 3e–3j). While population size increased along the eastern ledge of Northern Stellwagen Bank and in the northeastern corner of the northern box, nearly all other sectors saw a net loss in population size. The final size class (90 mm – ∞) exhibited growth along the coastline in the northwest corner of the northern box, the shallow zone of Northern Stellwagen Bank, and the deep zone bordering Northern Stellwagen Bank to the west. Additionally, the southern box exhibited a slight rise in population size within the shallowest area of Southern Stellwagen Bank, though surrounding change was mostly negative (Figure 3k). As dynamical phenomena (e.g., tidal movements, wind, storms) were considered in this run, aragonite saturation varied by location with time. Tidal movements increase mixing of waters, bringing highly saturated aragonite down to deeper waters and promoting growth. This downwelling effect was likely a critical driver of the increased growth within the deep area west of Northern Stellwagen Bank. Overall, full dependency on physics, aragonite, and temperature resulted in remarkable variation in growth within and among size classes.

Run 2, executed with a dependency on temperature and physics, exhibited results similar
to those of Run 3, where temperature was considered but aragonite dependency and physics were excluded. Slight differences were detected among plots, especially in Figures 4b and 5b (size class 45—50 mm) and Figures 4e and 5e (size class 60—65 mm). All other plots contained no visible variation. The lack of strong variation between runs displays that the influence of dynamical phenomena on temperature distribution in Massachusetts Bay has little impact on shellfish growth. However, there was a subtle effect of physics on population growth due to temperature over the 32.5-day period. Without physics, more change in population size was identified in shallower/warmer regions. When physics was considered and waters mixed, shallow water-specific growth became less notable. If the simulations were carried out over a longer span of time, this effect would be magnified.

Run 4 considered both temperature and aragonite saturation, but physical phenomena were excluded. Values for temperature and aragonite were read only from the initial conditions. As there was no opportunity for dynamical phenomena to influence the distribution of aragonite or temperature, the results of Run 4 shared little similarity with the results of Run 1. Run 4 saw considerable growth within areas of deeper water surrounding Northern Stellwagen Bank in size classes 6b-6j, but this growth was less frequent in the Run 1. Additionally, the last size class of the fourth run (Figure 6k) exhibited growth on the shallowest areas of Stellwagen Bank, whereas the last size class of Run 1 saw negative growth within these areas (Figure 3k). Dissimilarity between the two runs highlights the vital role that ocean dynamics play on aragonite and temperature distribution, as well as on shellfish growth and development.

4 Future Work

The 32.5-day run period of each experimental scenario provided valuable insight into the role physical phenomena and ocean acidification variables (aragonite saturation, tempera-
ture) play in influencing sea scallop growth. A longer run duration, perhaps over an entire season of growth, may reveal more significant trends. For instance, if the run time of Run 2 was extended, the effect of physical phenomena on temperature distribution may impact a wider range of size classes. An even longer time frame, utilizing data from before the industrial revolution, would convey the impact of anthropogenic ocean acidification over a more sweeping, historical timeframe. Runs observing these long time scales would prove critical in shaping an understanding of how shellfish growth rates are affected by long-term changes in ocean environments. The results generated by this model might also be compared to empirical data to test accuracy. Moreover, forecasting programs could be probabilistic, providing reliability estimates of produced forecasts. Additionally, initial conditions and recruitment parameters could be set as position-dependent, rather than evenly distributed, which would provide a precise profile of ideal growth zones. Such a system would provide invaluable advice to shellfish farms and fisheries management organizations regarding the threat of ocean acidification on a particular shellfish population.

Currently, the model only addresses the impacts of aragonite and temperature on larval survival (equation 2). Unlike their mature counterparts, sea scallop larvae are not fixed to the ocean floor. Instead, they are dispersed throughout the water column [41]. Larvae also have the ability to move and be advected horizontally. Modeled recruitment would be enhanced by functions and/or simulations representing the impact of aragonite and temperature on larvae across a wide depth and area range, rather than limiting these variables to deep-water values.

In Massachusetts Bay, several additional shellfish species exist. Of special relevance is the American oyster, which has substantial ecological and economic value [42]. Due to disease and improper management, American oyster populations have fallen significantly [43]. In fact, estimates conclude that harvest levels of the widely-studied Chesapeake Bay American oyster population have fallen to only 1% of maximal productivity [44, 45, 46]. The equations
and parameters implemented in the current sea scallop model could be adapted to this species, and a growth model could be developed to accurately predict the impact of ocean acidification on the strained American oyster population. Such an adaptation would involve evaluating the implications of shallow-water environments on the MIT-MSEAS PE model and considering the impacts of river input and erosion on coastal environments.

This study did not consider the effects of fishing pressure on modeled sea scallop populations, which assisted in reaching a baseline understanding of the impacts that aragonite, temperature, and dynamical phenomena have on sea scallops. To practically implement such a model into fisheries management, however, fishing mortality and selectivity for certain size classes should be incorporated.

5 Conclusion

When coupled to the MIT-MSEAS PE model, the shellfish growth model successfully projected the impact of aragonite, temperature, and physical phenomena on sea scallop growth in Massachusetts Bay. Trends exhibited in Run 1 displayed a distinct reliance upon aragonite by Massachusetts Bay sea scallops. Due to the wide range in growth rates within and among size classes, it can be concluded that Massachusetts Bay aragonite saturation was not evenly distributed. Temperature effects on sea scallop growth were less obvious (Run 2, Run 3) and appeared relatively uniform in distribution. When physical phenomena were incorporated, the model yielded subtle changes in shellfish growth. A temperature redistribution effect was noted when Run 2 was compared to Run 3. Similarly and more convincingly, when Run 1 and Run 4 are compared, the impact of ocean dynamics on aragonite distribution and scallop growth is demonstrated. Tides, internal waves, wind, and solitons are crucial drivers of this phenomenon, intermittently bringing highly-saturated surface aragonite down to deeper waters inhabited by sea scallops and promoting substantial growth and development. The
shellfish model produced by this study could be enhanced with longer run times and the consideration of fisheries mortality. Longer run times would model the centuries-long impact of ocean acidification on sea scallops, and the consideration of fisheries mortality would facilitate more relevant use by fisheries management organizations.

6 Practical Takeaways

This study’s findings have practical applications in the fields of aquaculture, fisheries management, and environmental assessment. The shellfish aquaculture industry is prominent along the Atlantic coast and in the state of Massachusetts, yielding 45.5 million dollars in a 2015 study from Augusto and Holmes [47]. This industry will require consistent growth to meet demand. As oceans acidify and shellfish growth rates slow, it will become harder to sustain the shellfish industry. With slight modification of initial parameters, the model in this study could be implemented by shellfish farmers along the coast. Farmers would be able to determine ideal locations and seasons for growth, increasing efficiency and providing the industry with more resilience against a changing climate.

While aquaculture has the ability to control early stages of growth in shellfish, the Atlantic coast mollusk fishery does not have this advantage. Ocean acidification poses an extreme risk to wild shellfish populations, and must be factored into stock assessments and regulation. The model produced by this study would provide valuable information regarding threatened areas and populations. To be useful to fisheries management, the model would have to be adapted to the specific location and goals of the fishery under study. Fishery selectivity of certain size classes or harvest locations could be factored into the model through mortality rates or population/length distribution. Already, models exist that consider the pressures of fishing on shellfish populations [21]. Different species and fishing areas have different regulations. These parameters could be fine-tuned to specific species and fisheries, generating an accurate
representation of the impact of ocean acidification on already-pressured shellfish populations.

Finally, the structure of this model and the program itself should be reviewed for use in environmental assessments. The implications of ocean acidification are not limited to sea scallops, nor to the Atlantic coast. This phenomenon impacts ecosystems across the world. The structure and function of this study’s model could be easily applied to a wide range of environmental conditions. Such application would increase the world’s understanding of the impact that a changing climate and acidifying oceans have on our marine ecosystems.

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Appendix

A.1 Shellfish Growth Model Pseudocode

Shellfish:

- Get_input_parameters
- Set_initial_conditions
- Loop_over_all_time
  * Read_data
  * Update_shellfish
  * Save_shellfish
- Close files

Get_input_parameters:

- Fetch initial parameters for growth model

Set_initial_conditions:

- Open MIT-MSEAS PE, output files for shellfish
- Input initial conditions (e.g., population distribution of shellfish, boundary parameters)

Read_data:

- Read salinity from MIT-MSEAS PE
- Read total alkalinity from MIT-MSEAS PE
- Read dissolved inorganic carbon from MIT-MSEAS PE
– Read temperature from MIT-MSEAS PE

Update_shellfish:

– Compute $G$
– Compute $R$
– Get $M$
– Total count of scallops in each size class calculated here

Save_shellfish:

– Save output

Compute_G:

– Get $H_\infty$
– Compute $H$
– Compute $K_{T,\Omega}$
– Growth matrix calculated here

Compute_R:

– Get $\alpha_R$
– Get $\gamma_R$
– Compute $scaled_{SSB}$
– Total recruitment calculated here

Compute $K_{T,\Omega}$
- Compute $K_i + \Delta K_t$
- Compute $\Delta K_\Omega$
- Total change in $K$ due to temperature and aragonite calculated here

Compute $scaled\_SSB$

- Compute $scale\_factor$
- Scaled SSB calculated here

Compute $K_i + \Delta K_t$

- Get $M_\infty$
- Get temperature from MIT-MSEAS PE
- Initial $K$ plus change in $K$ due to temperature calculated here

Compute $\Delta K_\Omega$

- Compute $\Delta G$
- Get $K_i$
- Change in $K$ due to change in aragonite calculated here

Compute $scale\_factor$

- Get average survival
- Get aragonite scale factor coefficient
- Get survival intercept
- Get aragonite from MIT-MSEAS PE
- Scale factor applied to spawning stock biomass calculated here
Compute $\Delta G$

- Get delta growth aragonite coefficient
- Get delta growth intercept
- Compute $\Delta \Omega$
- Relative change in growth calculated here

Compute $\Delta \Omega$

- Get aragonite from MIT-MSEAS PE
- Get $\Omega_{control}$
- Relative change in aragonite from the control value calculated here